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# Fractal Dimension Analysis Applied to Soil CO<sub>2</sub> Fluxes in Campotosto's Seismic Area, Central Italy

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**Abstract:** This article reports fractal dimension analysis applied to soil CO<sub>2</sub> fluxes measured in an Italian seismic area. The work was carried out with the use of a calibrated flux chamber unit. The fractal dimension (FD) from isotropic variograms was used as a method to understand related scale-dependent phenomena. The aim was to investigate the spatial variability of CO<sub>2</sub> flux measurements in four directions (horizontal, vertical, 45° and 135° directions) related to different distances between the measuring points and from a fault. High fractal dimension values were found ( $2.5 \le FD \le 3.0$ ). These imply strong anti-persistent behavior near to and far from the fault. Lower fractal dimensions were addressed at longer distances from the fault.

Keywords: fractal dimension; CO2 fluxes; seismic area

# 1. Introduction

Surveys of soil gas have been commonly used to trace buried faults and to study the behavior of endogenous gases of specific origins in the shallow environment (i.e., trace gases such as radon and helium, and carrier gases such as carbon dioxide, nitrogen, methane, etc.) [1–8] and have received significant attention over the last few years as earthquake precursors [6,9–15]. In addition, changes in stress/strain associated with seismic activity may force the migration of crustal fluid, in particular along active faults, thereby altering the geochemical characteristics of the surface fault zone [16–18]. The movement of these gasses by diffusion and/or advection along active faults will create flow anomalies, with concentrations substantially higher than the background levels; such anomalies may provide accurate information about the position and morphology of the shallow fracturing zone, as well as the permeability within the fault zone [3,4,7,19–21].

In recent years, geological, seismological, geophysical and geochemical knowledge has contributed to a much stronger, but far from complete, understanding of the link between faults and earthquakes in terms of space and time, with important advances in the analysis of seismic precursors applied to probabilistic seismic hazard assessment [12].

A fractal is a geometrically rough or broken structure that can be subdivided into sections, each of which is (at least approximately) a copy of the whole. Fractal geometry is the analysis of geometric



forms that appear chaotic or irregular when compared to those of standard geometries (e.g., line, square, sphere, etc.), which exhibit extreme orderliness. Fractals may show invariant properties under reasonable contractions or dilations [22].

The authors in [23,24] asserted that phenomena with scale-dependent spatial variability can be studied through the concept of the fractal dimension. The technique has also been applied to non-continuous spatial and temporal phenomena [25]. According to [26], the fractal dimension applied to the characterization of soil can provide an estimation of the scale regularity and irregular behaviour. Scale dependency and spatial variability have been addressed in the relationship between CO<sub>2</sub> fluxes and soil attributes [27,28]. In addition, [29] reported that the spatial variability of CO<sub>2</sub> flux is partially subject to experimental semi-variogram adjustments, which must be properly selected. This subjectivity can be attributed to the dependence of the experimental semi-variogram on grid characteristics, such as the direction and sampling distance used at the experimental site [30,31]. Previous studies have used different range values of CO<sub>2</sub> fluxes for different locations, soil types and vegetation covers [32–35]. Certainly, new approaches and more research are needed to better understand the spatial variability of CO<sub>2</sub> fluxes at different scales [36]. Some studies were carried out to understand the fractal behavior in seismic areas [37,38]. According to [39,40], the CO<sub>2</sub> fluxes in seismic areas can be used to estimate the relationships between CO<sub>2</sub> gas emissions and seismic activity.

Accounting for the above findings, this paper focused on the study of the spatial analysis and variability of  $CO_2$  flux from soil using fractal dimensions calculated with the semi-variogram method. The fractal characteristics of the underlying mechanisms that govern the spatial variability of  $CO_2$  flux were delineated, and important information about the hidden long-lasting patterns of the Campotosto seismic area (Central Italy) were found and reported.

#### 2. Methodology and Data Analysis

#### 2.1. Experimental

The study area is located at Campotosto ( $42^{\circ}55'$  N and  $13^{\circ}38'$  E), a small village approximately 20 km away from L'Aquila city, in the Abruzzo region (Central Italy, Figure 1). Abruzzo lies in the Apennine chain. In the vicinity of Campotosto, there is a significant active fault, the so called Mount Gorzano fault, of approximately 28 km length (Figure 2), a normal-type fault with WSW-dipping and a length-in-depth about 19.6 km [41–44]. Moreover, the rocks that outcrop in this study area are a flysh-type succession called the pre-evaporitic member of the Laga Formation [45]. This specific stratigraphic succession found in the study area is the arenaceous-pelitic (LAG4, lower Messinian [46]), characterized by the alternation of layers of sandstone and clay with relation A/P > 1 [46]. This is a strategical study area because is very near to active faults that generated big earthquakes on 24 August 2016 and also on 26 and 30 October 2016 in Central Italy [43,44].

The 37  $CO_2$  flux measurements were performed, independently, in two discrete phases, on 25 September 2019 and on 1 December 2019. Figure 3 presents all these 37 different measurements in the Campotosto area. During all the measurements, the climatic conditions were sunny without previous rain. Prior to data acquisition, superficial vegetation was removed in order to achieve the direct contact of the measuring device with the soil and rock underneath.

CO<sub>2</sub> fluxes from soil were measured by a flux chamber unit constructed at the University 'G. D'Annunzio' of Chieti-Pescara, Italy. The instrument consists of a cylindrical stainless steel container of 30 cm diameter and of 18 cm height (12.72 L volume-chamber). The unit was calibrated and tested for measurement accuracy. The calibration phase was implemented through the use of a laboratory-constructed system, which controlled the flow of various air and CO<sub>2</sub> mixtures (via a LabView program (National Instruments Corporation, Austin, TX, USA)), while concurrently measuring the flow with an external Mass Flow Controllers made by MKS Instrument, type 1179 Mass-Flo<sup>®</sup> Controller (Wilmington, MA, USA) and an internal CDM4160 (Figaro<sup>®</sup>) pre-calibrated module that was installed inside the flux chamber for carbon dioxide measurement (value range: 390–4000 ppm of CO<sub>2</sub> concentration).

The calibration factor of the flux chamber unit was calculated as the slope of the least-squares fit between the  $CO_2$  concentrations measured by the chamber unit (internal controller) and the flow-driven  $CO_2$  concentrations measured by the external units.



**Figure 1.** Map of study area in proximity to Campotosto in Central Italy. The figure presents the Mount Gorzano fault (red line) [43,44] and the main cities.



Figure 2. Geomorphological map of the study area. Map taken from [45].



**Figure 3.** The map of the full set of measurements (flux values in ppm cm<sup>-2</sup> s<sup>-1</sup>), with the fault position [43,44] shown with a red line. The x axis represents the latitude from  $42^{\circ}33'23.436''$  N to  $42^{\circ}33'38.286''$  N, while the y axis represents the longitude from  $13^{\circ}22'58.962''$  E to  $13^{\circ}23'17.3832''$  E.

All the fluxes were calculated according to [47] by the equation:

$$\phi_{\rm gas} = \alpha H_c \tag{1}$$

where  $\alpha$  is the slope of the increase in gas concentration over time and  $H_c$  is the height of the chamber. Using Equation (1), the CO<sub>2</sub> fluxes were evaluated, and the relative results are shown in Figure 3 in ppm cm<sup>-2</sup>s<sup>-1</sup>.

The test phase was implemented to determine the quality of the  $CO_2$  concentration measurements provided by the internal module of the unit. Two main tests were carried out: the first was implemented through increasing steps of the  $CO_2$  concentrations driven to the chamber unit, and the second, through decreasing steps of  $CO_2$  concentrations. Between each experimentation, adequate time was left for the  $CO_2$  concentration values of the chamber to return to the corresponding baseline ones, which are usually around 400 ppm. In this way, the dead-time interval of the flux chamber unit was calculated, i.e., the time needed for the unit to return to its baseline recordings.

Proper sensors inside the chamber measured the temperature, pressure and humidity. Specifically, the LM35CAZ Precision Centigrade Temperature Sensor (National Semiconductor<sup>®</sup>, Santa Clara, CA, USA) was employed for temperature measurements, the XFPM-115KPA (Fujikura<sup>®</sup>, Tokyo, Japan) was used to measure atmospheric pressure, and the Honeywell<sup>®</sup> (Charlotte, NC, USA) HIH-4000 Series was utilized for relative humidity (RH) measurement. Moreover, the pressure inside and outside the chamber was balanced by a pressure compensation valve located at the top of the chamber. A rotating rod of carbon fibre and a gear motor (set to rotate less than 60 r/min) located inside the chamber reduces any bias effect due to gas stratification inside the chamber. The data were stored and saved every second with an Arduino board, mounted at the top of the chamber.

#### 2.2. Mathematical Aspects

Spatial analysis based on the direction of measurements may provide important information regarding the fractal dimension (FD) and related scale-dependent phenomena [24,29,36]. Using experimental variogram analysis [36,48], it is possible to determine the fractal dimension according to Equation (2).

$$\hat{y}(h) = \frac{1}{2}N(h)\sum_{i=1}^{N(h)} [Z(x_i) - Z(x_i + h)]^2$$
<sup>(2)</sup>

where  $\hat{y}(h)$  is the semi-variance of points separated at a distance *h*, *N* is the number of pairs of points separated by *h* distance,  $Z(x_i)$  is the value of the attribute at position  $x_i$  and  $Z(x_i+h)$  is the value of the same attribute in position  $x_i+h$ .

The spatial structure of fractal surfaces can be described by means of a power law, as shown in the following relation:

$$|z(x) - z(x+h)| \propto h^H \tag{3}$$

where the location *x* has an attribute value *z*, with the separation distance *h*, and *H* is the fractal codimension or Hurst exponent [49]. If  $0 < H \le 1$ , the fractal codimension can be defined as follows:

$$H = d - FD \tag{4}$$

where FD is the fractal dimension and d is the Euclidian dimension of the system in which the fractal distribution has been described. For lines, surfaces and volumes, d = 1, 2, 3, respectively. Hence, for a distribution of a soil attribute, the fractal dimension is given by

$$FD = 3 - H.$$
 (5)

Comparing Equations (2) and (3), it is possible derive the semi-variance at a given scale as

$$\hat{y}(h) \propto h^{2H} \tag{6}$$

or as:

$$\log[\hat{y}(h)] \propto 2H \log[h]. \tag{7}$$

Through a logaritmic transformation, the slope of the experimental variogram can be calculated on the log–log scale as 2*H*. The *H* exponent can then be obtained by means of a linear regression on the log–log graph by the following equation [23]:

1

$$\left(\frac{\log[\hat{y}(h)]}{2\log[h]}\right).$$
(8)

When H = 0, the value of FD is equal to 3, which represents the lack of a spatial variability structure, as there would be no relationship between the way the attribute varies in space and the distance between the points. In this case, there is no fractal dimension, and the methodology does not apply. However, when 0 < H < 1, the fractal dimension has values that characterize the presence of spatial variability and dependence with h [29,31]. According to Equation (5),  $2.0 \le FD \le 3.0$ . When  $2.0 \le FD \le 2.5$ , there is a persistence behavior, while if  $2.5 \le FD \le 3.0$ , there is an anti-persistence behavior [29,36].

#### 3. Results

The results of the calibration phase are given first. These results are very important because they verify the CO<sub>2</sub> measurements provided by the flux chamber unit and validate the method of the measurements. As aforementioned, two calibration factors were calculated, corresponding to increasing and decreasing steps of CO<sub>2</sub> driven concentrations. These are characteristically shown in Figure 4. The least-squares fit for Figure 4a is y = 1.1008x + 1.4180 with  $R^2 = 0.9759$ , and for Figure 4b, y = 1.0075x + 2.2513 with  $R^2 = 0.9771$ . Both lines have a very high square of the Spearman's correlation coefficient. This means that the flux chamber unit very accurately records the pre-determined CO<sub>2</sub> concentrations driven inside and originating from the external mass flow controllers. It is worth mentioning that the two mass flow controllers are good devices and that, therefore, the pre-determined driven CO<sub>2</sub> concentrations are very well defined. Hence, it is expected that in real-time situations, the CO<sub>2</sub> measurements will be accurate and reliable.



**Figure 4.** The best-line fits of the two tests: the first, (**a**) reports the increasing concentration test, while the second, (**b**), the decreasing concentration test. In both figures, the x axis (ppm  $CO_2$  (d)) reports the  $CO_2$  concentrations driven into the flux chamber unit, while the y axis (ppm  $CO_2$  (m)) reports the  $CO_2$  concentrations measured.

The measurements carried out in the Campotosto area to delineate the CO<sub>2</sub> flux profiles of the Mount Gorzano fault and its vicinity are shown in Figure 3 The map shows the 37 CO<sub>2</sub> flux measurements that were performed at various distances and positions from the fault. The maximum flux measured was 25.84 ppm cm<sup>-2</sup> s<sup>-1</sup>, and the minimum one was 2.858 ppm cm<sup>-2</sup> s<sup>-1</sup>. As can be

observed from Figure 3, high CO<sub>2</sub> flux values (from 20 to 25 ppm cm<sup>-2</sup> s<sup>-1</sup>) are found near the fault, while, on the contrary, lower flux values (from 2 to 4 ppm cm<sup>-2</sup> s<sup>-1</sup>) can be observed in positions far from it. Tectonic fracture is related to the deformation pattern along the fault system and also facilitates a pathway for the  $CO_2$  gasses [50–52], therefore leading to elevated fluxes near the fault. It should be noted that the soil  $CO_2$  flux can be influenced by the rock type and by other parameters like chemical, physical and biological processes, including the soil temperature and atmospherical pressure [28,53]. Since the confirmation of the above observations needs further analysis, it is crucial that appropriate methods are employed. A significant approach towards this is monitoring the spatial fractal behavior of the  $CO_2$  fluxes [29,36]. In association, it is significant to mention that the measurements were conducted following, as much as possible, the perpendicular direction (45° direction) with respect to the fault position. The scope was to acquire adequate profiles for statistical analyses. There are a variety of heights and spatial distributions of the measuring points in this study. The different distribution of heights is due to the morphology of the study area. Nevertheless, it seems that the surface profile of the fault is not a straight line (see Figure 3). Since the experimental approach of 45 degrees could not be followed exactly due to the morphology, further analysis in terms of the fractal semi-variogram method was attempted. In the following sections, this analysis is presented in detail.

As mentioned in the methods section and according to [36,54,55], the semi-variogram method is very efficient for the delineation of the hidden spatial trends in the CO<sub>2</sub> flux measurements. Since a spatial distribution of points has been achieved from the measurement procedure, the following alternative methods were employed for the calculation of fractal dimensions with the semi-variogram method: (1) reorganizing the flux measurements in four different directions (horizontal, vertical,  $45^{\circ}$  and  $135^{\circ}$ ) via arbitrarily selecting pairs of measurements from the whole dataset (each pair indicated with a different ID number), so that all the selected series of pairs correspond to each one of the four directions; (2) reorganizing the series of pairs of 1), with reference to their intra-pair distance; and (3) arbitrary selecting pairs of measurements with different distances from the fault. The fractal dimensions with reference to the horizontal direction (Figure 5a) exhibited values of  $1.5 \le \text{FD} \le 3.0$ . The fractal dimensions with reference to the  $45^{\circ}$  direction (Figure 5c) all had values over 2.5; in particular, certain fractal dimension values were found of  $2.85 \le \text{FD} \le 2.9$  and  $2.8 \le \text{FD} \le 2.85$ . The fractal dimensions with reference to the  $135^{\circ}$  direction (Figure 5d) present values of  $2.0 \le \text{FD} \le 3.0$ .

Figure 6 presents fractal dimensions with reference to the distance in the four spatial directions. These are the distances between the various points that are spatially dispersed. Due to the scattering of the measurement points, there are a variety of distances from 20 m to 140 m. Note that Figure 6 is a reorganization of Figure 5 with reference to the distance and not to the ID number. Since there is no preferred direction in the above distances, Figure 6 has a different distribution of values in both the horizontal and vertical axes. It is very important to mention that all the fractal dimension values are over 2.5. This is a very significant finding that implies a strong fractal behavior of the  $CO_2$ -generating system in the spatial manner because the Hurst exponent (H) is between 0 and 0.5 [47,55]. This significant finding implies that when a high CO<sub>2</sub> concentration value is found, the next value will be low, and this tendency follows throughout the Campotosto area. Strong anti-persistent, long-lasting interaction exists. Observing that most high values are near the fault, the anti-persistence behavior justifies that the values after the fault will be low while the remaining values will have a trend of higher–lower values; all these values are lower than the corresponding values on the fault [56]. The fractal dimensions in the horizontal direction (Figure 6a) were all over 2.5; in particular,  $FD \ge 2.9$ at 40 and 80 m;  $2.8 \le FD \le 2.9$  at 50, 100 and 120 m;  $2.6 \le FD \le 2.7$  at 30, 90 and 110 m; and FD  $\le 2.6$  at 60 m. The fractal dimensions in the vertical direction (Figure 6b) are all above 2.5; in particular, the FD is above or equal to 2.9 at 40 and 90 m;  $2.8 \le FD \le 2.9$  at 60 m;  $2.7 \le FD \le 2.8$  at 70, 90, 100 and 130 m; and FD  $\leq$  2.7 at 140 m. The fractal dimensions in the 45° plane (Figure 6c) are all over 2.5; in particular,



 $2.7 \le FD \le 2.8$  at 20, 30 and at 50 m; and  $2.6 \le FD \le 2.7$  at 40 and 60 m. The fractal dimensions in the 135° plane (Figure 6d) are between 2.90 and 2.98.

**Figure 5.** Fractal dimension (FD) values found in different planes in four spatial directions. (**a**) FD values in the horizontal plane. (**b**) FD values in the vertical plane. (**c**) FD values in the 45° plane. (**d**) FD values in the 135° plane.



**Figure 6.** Fractal dimension (FD) values with different distances in the four spatial directions. (a) FD values in the horizontal direction. (b) FD values in the vertical direction. (c) FD values in the 45° direction. (d) FD values in the 135° direction.

This is characteristically shown in Figure 7 where the fractal dimensions with reference to the distance from the fault are presented. Figure 7 is the result of all the fractal dimension values taken at different positions from the fault. It is evident that the fractal dimension values away from the fault are lower. This finding verified the above facts in an alternative way. The fractal dimension values at greater distances are lower and, therefore, they follow the tendency of the CO<sub>2</sub> concentrations to exhibit anti-persistent behavior of lower intensity at greater distances. This means that the system goes away from strong anti-persistence and fractality when it is away from the fault; strong anti-persistence gradually diminishes. The significance of the strong anti-persistence behavior has been mentioned by several publications [56–60]. In the above publications, it was reported, in a multi-facet way, that the strong anti-persistent behavior, similarly to the strong persistence behavior, is a footprint of hidden activity in soil radon or environmental electromagnetic disturbances. According to [12], the radon in soil and  $CO_2$  fluxes from soil may have similar origins inside the Earth's crust prior to seismic activity. It is the micro-crack branching, as possibly in the Mount Gorzano fault, that may act as a pathway for the increase in radon and  $CO_2$  concentrations from soil. According to [61–63], the  $CO_2$  fluxes increase inside the fault. All these aspects report different faces of the same aspect; the Mount Gorzano fault, as a seismic area, is associated with a strong CO<sub>2</sub> pathways and increased CO<sub>2</sub> concentration fluxes.



Figure 7. Fractal dimension values from the fault (at distance 0) to the different distances.

The above findings are verified by the results of other investigators. For example, [29] reported similar results showing comparable fractal dimension values between 2.5 and 3.0 of the spatial variability of soil CO<sub>2</sub> emissions, where it was shown that the fractal dimension decreased, as well, with the distance. This very important finding is similar to the tendency shown in Figure 7 and should be emphasized. This tendency of a decrease in the CO<sub>2</sub> values with the distance was also reported by these authors. This is an additional verification of the findings of this work. Additionally, [36] reported a similar range of fractal dimension values as in this work, where the fractal dimension values were between 2.5 and 3.0 from 5 m to 60 m, also decreasing with the distance. Furthermore, [55] calculated—also with the experimental variogram method, using the same equation as Equation (2) of this publication—the spatial variability of soil CO<sub>2</sub> fluxes. All the findings of this paper are comparable to those of other related papers, despite the fact that the sample size was not very high. More or less, the accuracy of the employed measuring device in conjunction with the modern methodological approach contributed towards this achievement.

## 4. Conclusions

The technique used for this work was an experimental semi-variogram that was applied to measuring  $CO_2$  fluxes at the Campotosto seismic area (central Italy). Non-continuous spatial and temporal phenomena were addressed from the measurements.

The fractal dimension values at greater intra-distances and greater distances from the fault are lower and, therefore, they follow the tendencies of the  $CO_2$  concentrations, to exhibit anti-persistent behavior of lower intensity at greater distances.

The fractal dimensions of this work are within the value ranges reported in similar work.

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